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Determining the Mean Level* Discharges of Water in any Line of
Direction of Small Rivers

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DETERMINING THE MEAN LEVEL *
DISCHARGES OF WATER
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G. N. Petrov

(Submitted by Academician A. I. Nekrasov 23 January 1950)

To determine the mean level* discharges of small rivers
 (1)

In the absence or inadequacy of hydrometric data, N. A. Velikanov proposed a solution, based on the application of the methods of 3-dimensional interpolation and hydrological analogy, taking the water genesis into account, which, according to L. Lichkov (2), is determined by the geological structure of a locality. The mechanical application of these methods, in the form of isohyetal charts representing the modulus of the annual water run-off, does not give any satisfactory results (3), since the analogy was established without first ascertaining the homogeneity of the water genesis, and the interpolation was made without due investigation of the continuity of physico-geographical conditions.

It is known that stable mean level discharges are determined by underground feed waters. Many methods were proposed, therefore, for the determination of the values of the underground feed.

Hydrologically speaking, underground feeding is determined distinguishing by the separation, in the hydrograph, the stable portion according to minimum discharge value, and the variable portion ~~as~~ the difference between the high stable discharge value and minimum discharge value, after cutting off the rain floods values. However, this method does not clarify the distribution of underground feeding in a river system, causes and effects and does not discover the causality and stipulated interdependency of this phenomenon. Moreover, a minimum discharge value, according to this

(* Note: The word "level" as used here and in the text stands for the special technical phrase "mean level of water in a river or lake".)

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the given by the method, is determined from curve equation $Q = \varphi(H)$, with an error up to

100 - 300 per cent, due to permanent deformations in the river bed, ~~unsteady operation~~ ^{power plants} of water mills and hydro-electric ~~bottom~~

~~area~~, sharp temperature drop (in the winter), evaporation (in the summer). However, ~~over~~ even a perennial period of observations, this method does not always ~~determine~~ ^{ever} correctly ~~the~~ underground

~~super~~ feeding of the ~~river~~ basin. For instance (see Figure 1),

in the Tigriver, the summer ~~water~~ ^{of water} discharge may be diminished toward the estuary on account of transpiration from the bottom ~~area~~.

Penetration and infiltration into the river basins.

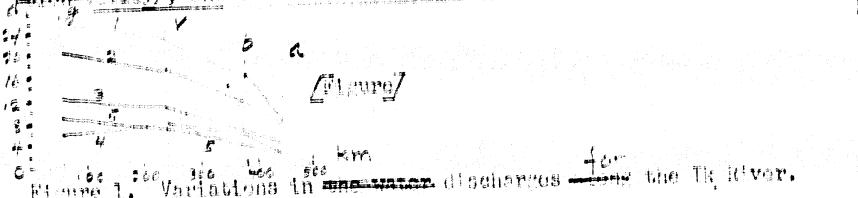


Figure 1. Variations in ~~river~~ discharges ^{for} the Tigriver.

measured in 1968 by the hydrological expedition of the Kazan' Institute, Academy of Sciences USSR. 2 - Mean perennial for August; 3 - mean perennial for January; 4 - minimum monthly mean for the summer; 5 - minimum monthly mean for ~~the~~ winter. a - observation point Kulabay; b - observation point Apsalyamovo; v - observation point Hayybak; g - observation point Podgornyye Raylyary; d - estuary.

Underground feeding can be determined more correctly (4) by the F. A. Nakarenko and B. I. Kudelin (5) methods, which require detailed hydrogeological investigations. These methods cannot be applied to the numerous small rivers, since the ~~point~~ between the river waters and the underground waters is diverse and cannot be established by a few supporting sources or typical lines of direction.

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When it becomes necessary to obtain data for numerous lines of direction of any small river, the magnitude and distribution of underground feeding along the river system can be determined hydro-metrically, with sufficient ~~simplification~~ ^{accuracy} and precision, by A. A. Trufanov's⁽⁷⁾ method, as the difference ~~in~~ ⁱⁿ the discharges of adjacent lines of direction. The ~~coefficient~~^{water-bearing} capacity of small rivers of stable ~~flow~~^{level*} is an objective index of underground feeding, and the study of the variations of ~~discharge~~^{level*} along the ~~course~~^{length} of a river, clarifies the part played by physico-geographical conditions in the formation of the ~~boundary~~^{canal} discharge and the genesis of the waters. Investigations of rivers (8-10) consist of measuring ~~and~~ great quantities of water discharged during the period of summer stable ~~flow~~^{level*} along the ~~course~~^{length} of the entire river net, and ~~in~~ ascertaining ~~the~~ the potential discharge based on the analysis of the physico-geographical factors. The effect of the latter cannot be established a priori, nor ^{can} hydrological zoning (11) be done merely on the basis of coincidence of individual geographical features (see Table 1), or territorial proximity of rivers. [Note: Table 1 is in the appendix.]

~~The~~ changes in the lithological composition and ~~the~~ thickness of the stratigraphic horizons, the presence of a deluvial-alluvial speleogenic covering, fissures, cavern processes, and, particularly, ~~the~~ tectonic structure, all go into the evolution of the peculiar local conditions for the movement and drainage of underground waters, which have a sharply pronounced effect on the mean ~~water~~ level of small rivers. (12)

Short-period Transitory investigations of rivers (9) give a picture of, merely ascertain the

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aquiferous capacity of a river, ~~in conformity with~~ corresponding to the climatic conditions of a given year. Their results cannot be utilized for ~~purposes of~~ computation of typical mean values, since the stability of mean level ~~and~~ fluctuations, even for river basins lying within a single climatic zone, is not the same and ~~depends~~ upon a very complex combination of physico-geographical and, mainly, biological conditions (see Table 2).

Table 2. Ratios of mean August ~~mean~~ discharge for the period 1932-1946, (in liters per second from basin area of one square kilometer)

River and point of observation	Measuring year	Low-water year	Median ratio of between high-water and low-water years
Raznitsa, R. Verbyzoidi	2.72	1.27	2.1
Ik, Narynsak	3.06	0.67	4.5
Ik, Kulbay	2.20	0.70	3.1
Ik, Lake Lebedinnoye	3.65	0.89	4.2
Sviyaga, Tyryshyevka	2.34	0.52	4.4
Sviyaga, Burunduki	1.15	0.23	5.0

Figure 7 [see page 7 for the figure]

Figure 2. Variation in ~~in~~ water discharges for typical Ik River tributaries; 1 - discharge measured during investigations in 1948; 2 - ~~sum~~ total influx ~~of tributaries~~ during 1948; 3 - Probable mean-annual discharge; 4 - Probable mean-winter discharge; 5 - Basin area in square kilometers; 6 - the cover line P_2^{Kaz} in absolute markings, A - Rya River; B - Killya River; V - Dymka River.

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It is not possible to investigate all rivers during all the years of their varying water cycles [high-water years and low-water years], in order to ascertain the mean ~~water~~ discharges along their course. Investigations have shown that to accomplish the above it is only necessary to ~~select~~ take a single hydrological "photograph" of the mean ~~water~~ level. To ~~determine~~ obtain the probable mean ~~amount~~ and ~~greater~~ even winter ~~water~~ discharge, it is necessary to subject the results of ~~the~~ short-time hydrological investigations to a hydrogeological purpose of clarifying the possible variations in quantity, for the ~~estimation~~ of the ~~future~~ ~~water~~ discharges under ~~various~~ ~~existing~~ ~~conditions~~, ~~especially~~ for the future ~~possibilities~~. The correctness of these assumptions and computations along the ~~length~~ of ~~the~~ river is ~~assured~~ indicated by the proportionality of water yield and hydroelectric power plants, and in comparison with the results ~~with~~ the ~~base~~ line of direction, ~~of~~ ~~the~~ ~~river~~.

~~For~~ ~~example~~, In the absence of a line of direction based on permanent observations of the river under study, the method of hydrogeological analogy is applied.

This method was used for determining the mean August and January ~~water~~ discharges along the ~~length~~ of the following river systems: Ily, Syriyga, Shesima, Zay, and others (See Figure 2).

The Rya River lies in a ~~valley~~-like tectonic depression. Considerable and conformable declivities on the terrestrial surface and in subterranean relief are conducive to the rapid run-off of surface and underground waters, particularly since the river system is cutting through the strata of sandstone (P_2^{Kazbel}) having variable aquiferous horizons. The stability of ~~water~~ discharges in magnitude ~~over~~ time is insignificant. The ~~direction~~ of the reduction

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in measured discharges ~~based on the~~ is based on the
~~calculated~~ discharges ~~based on the~~ work
 discharges of previous years and the ~~performance~~ of water mills.

which are devoid of
 The Killiva river cuts through geological formations ~~with~~
 aquiferous horizons, (P_1 , K_2 , S_2 , I_2) The discharge of the
 river system is ~~based~~ at the headwaters of the rivers and its
 tributaries from the upper horizons (P_1 , K_2). In consequence of
 the great variability of depositions, there are variations in the modulus
 of discharge; but ~~the~~ the stratification ~~ensures~~ [shallow dip] a
 high degree of regularity of underground-water and stability
 of water discharge. The effect of
 the magnitude of elevation of the alluvial deposits accounts
 for the ~~increased~~ ~~discharge~~ in the river-bed of the Killiva
 more, no account of the influence of the tributaries.

Along the course of the Lyman River there are sections having
 zero, positive, or negative ~~discharge~~ due to the various dips
 in the rock formations, although the river bed ~~which~~ is
 situated in stratigraphic horizons ~~similar~~ to those of the Killiva
 river. Individual tributaries and river sections, having ~~considerable~~
 storage ~~capacity~~ in comparison with that on the
 underground water ~~capacity~~ in the ~~stratigraphic~~ surface, with ~~considerable~~ ~~variable~~ dip ~~capacity~~ stratifi-
 cation, insure the ~~regularity~~ of the water discharge. The discontinu-
 ity of the water discharge at the estuary is due to the ~~effect~~ of (influence
 of caverns) ~~on~~ (13).

The application of ~~this~~ utilizing the factual data
 insures the most ~~accurate~~ determination of the probable mean level
 discharges ~~discharge~~ in any line of direction of a river system, which
 is ~~necessary~~ of great importance under the existing conditions

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~~spread~~
 of wide utilization of small rivers. Precision in the computation
 of the rate of discharge of ~~the~~ water is determined by the degree
 of knowledge gained through research and investigation of an area,
 correctness and
 and the completeness of the analysis of the prevailing physico-
 geographical conditions.

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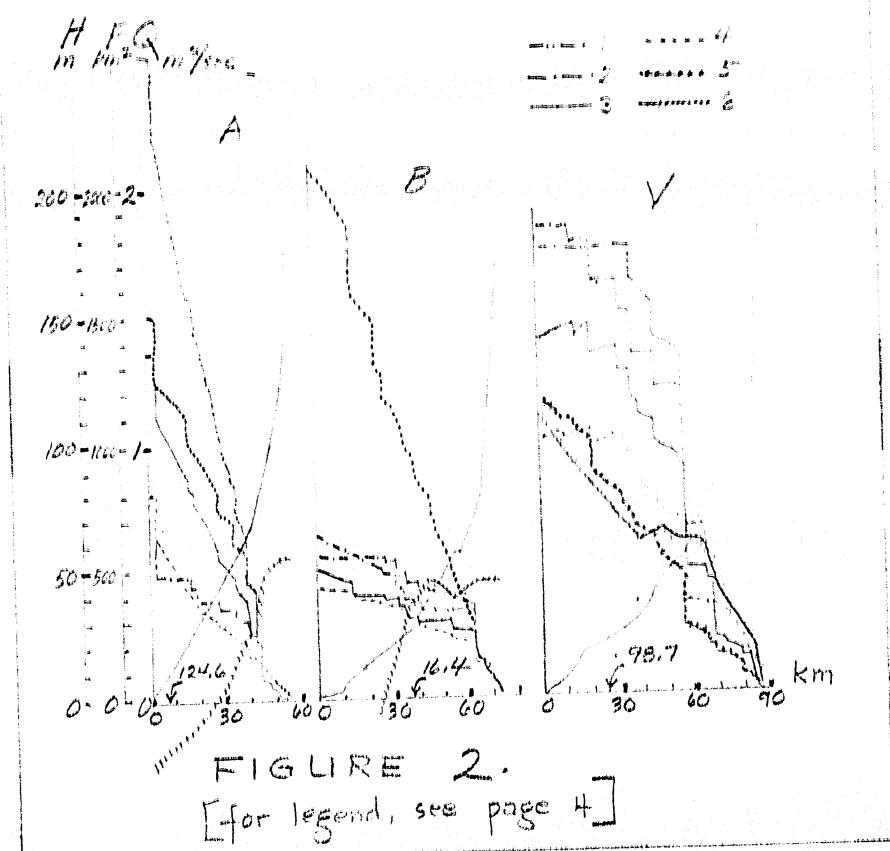
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